

Reports

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Critical Loads of Acidity for

High Precipitation Areas

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Project Coordinator: Helmut Hojesky

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CRITICAL LOADS OF ACIDITY FOR HIGH PRECIPITATION AREAS

Results from a Workshop held in Vienna, March 9-10 1992 on the Modifications of Austrian, German and Swiss Maps of Critical Loads of Acidity.

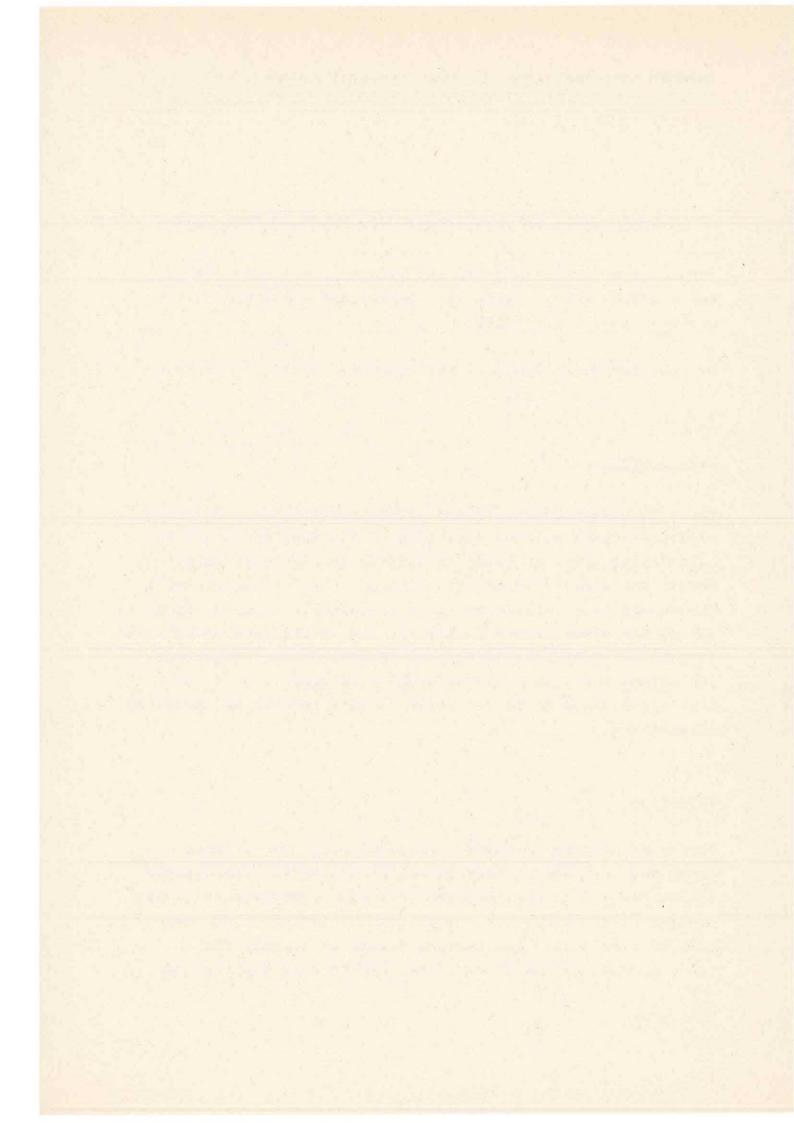
by Umweltbundesamt/Federal Environmental Agency (Austria)

Introduction

There have been some problems using the originally proposed simplified mass balance equations of the Mapping Manual for calculating critical loads of acidity in high precipitation areas. The simplified equation appears to be biased towards the runoff rate. This results in unrealistic critical loads values for alpine areas in Austria and Switzerland as well as in the Harz mountains of Germany and the German Erzgebirge. The values obtained with the simplified equation indicate, that areas known to be sensitive to acid deposition appear as insensitive.

Objective

The objective was to modify the simplified mass balance equations, to remove simplifications not valid, and thereby remove the runoff bias observed for high precipitation areas. The modified versions of the critical loads equations were used to reculculate the critical loads of acidity for Austria, Germany and Switzerland, and to draw modified maps.

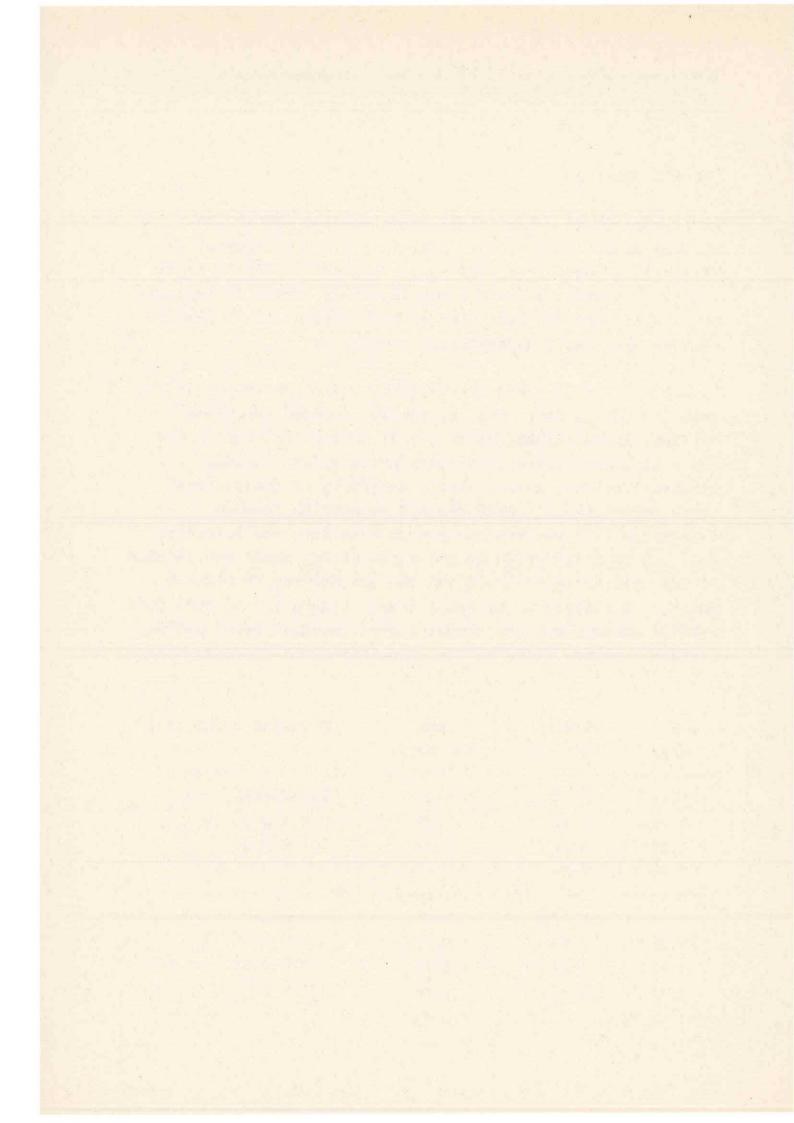


The old equation

As it appears in the critical loads mapping manual and related documentation, the originally proposed simplified mass balance equations (SMB) were derived on a basis of a series of assumptions and simplifications. These assumptions and simplifications have been worked through and checked, whether they apply under alpine conditions.

In Austria, Germany and Switzerland, where the runoff varies from 0,3 to 3,5 m/yr, the originally proposed SMB gives critical loads values, which are obviously too high. Areas known to be sensitive to acidification appear as being insensitive. That means, that - according to the critical loads calculations - they are are allowed to receive significantly more acid deposition than they can tolerate. This can be illustrated by the table below, where the results of the originally proposed SMB and the results of PROFILE, which is a complex multi-layer model working on the same mass balance principles, are compared for 3 Swedish sites using different runoff rates in the calculation of critical loads:

Q	PROFILE	SMB	Site and input data
m/yr	keq/ha yr	keq/ha yr	
0.15	0.45	0.67	Sandvatten:
0.55	0.41	1.03	W = 0.15
1.15	0.21	1.57	D = 0.6
2.15	0.20	2.47	U = 0.5
0.15	1.00	1.80	Gårdsjön:
0.55	1.06	2.30	W = 0.68
1.15	1.20	2.92	D = 0.55
2.15	1.06	3.90	U = 0.5



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W = weath	ering rate	e of Ca + Mg +	K + Na [ed	[/(ha.yr)]	
 2.15	3.10	6.74	U =	0.7	-
1.15	2.75	6.00		0.75	
0.55	2.45	5.30	W =	2.0	
0.15	2.25	4.95	Skåne	no. 7:	

D = base cation deposition [eq/(ha.yr)]

U = base cation uptake [eq/(ha.yr)]

The formula applied in the originally proposed SMB was

$$W + D - U$$

$$CL = W + 0.09 \cdot Q + 1.5 \cdot (-----) \qquad (1)$$

$$(BC/Al)_{crit}$$

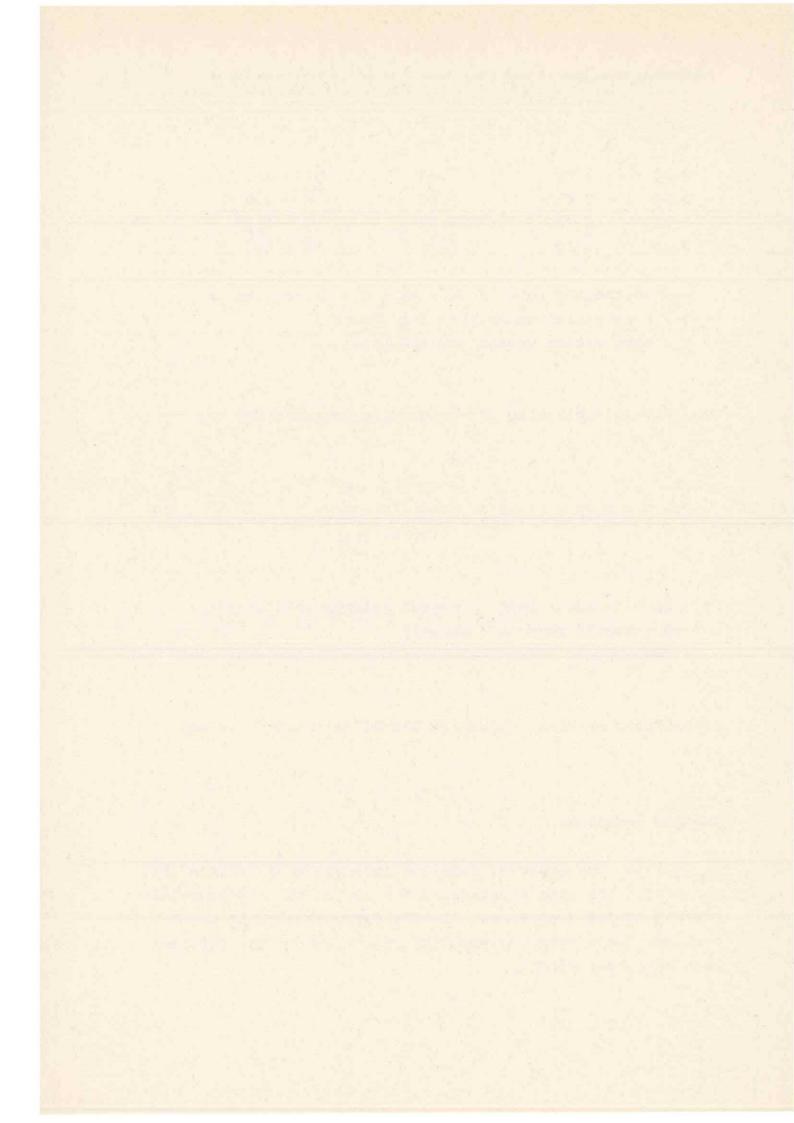
CL = critical load of actual acidity [eq/(ha.yr)] Q = runoff rate [m³/(ha.yr)] (BC/Al)_{crit} = critical base cation : aluminium rate

As critical chemical criterion the BC/Al-ratio = 1,0 was used.

The new equation

By deriving the equation from the complete mass balance, but not making the same simplifications as for the old equation, one can derive a different formula for the calculation of critical loads starting with the definition of the critical load of actual acidity:

3



$$CL = W - ANC_{l(acc)}$$

ANC = acid neutralizing capacity = alkalinity $[eq/m^3]$ ANC_{1(acc)} = acceptable alkalinity leaching [eq/(ha.yr)]

In this equation the limiting ANC leaching is determined by the maximum permitted leaching of H^+ and AL^{3+} :

$$ANC_{l(acc)} = -H^{+}_{limit} - Al^{3+}_{limit}$$
(3)

The limiting Al-flux is determined by the molar BC/Al-ratio applied, which depends on a mass balance of available base cations. The simplified SMB had assumed, that all the production of alkalinity due to weathering is equivalent to the base cation production involved in the BC/Al-ratio. This is different from nature, because approximately 20-40% or more of the base cations produced by weathering is Na, providing no protection against Al and thus not participating in the BC/Al-ratio:

$$W(Ca+Mg+K) = (1-X_{Na}) . W(Ca+Mg+K+Na)$$
(4)

1 - $X_{Na} = X_{BC}$ X_{BC} = Fraction of weathering excluding Na = 0,6-0,8

This leads to:

$$BC \qquad (1-X_{Na}) \cdot W + D - U - L_{min}$$

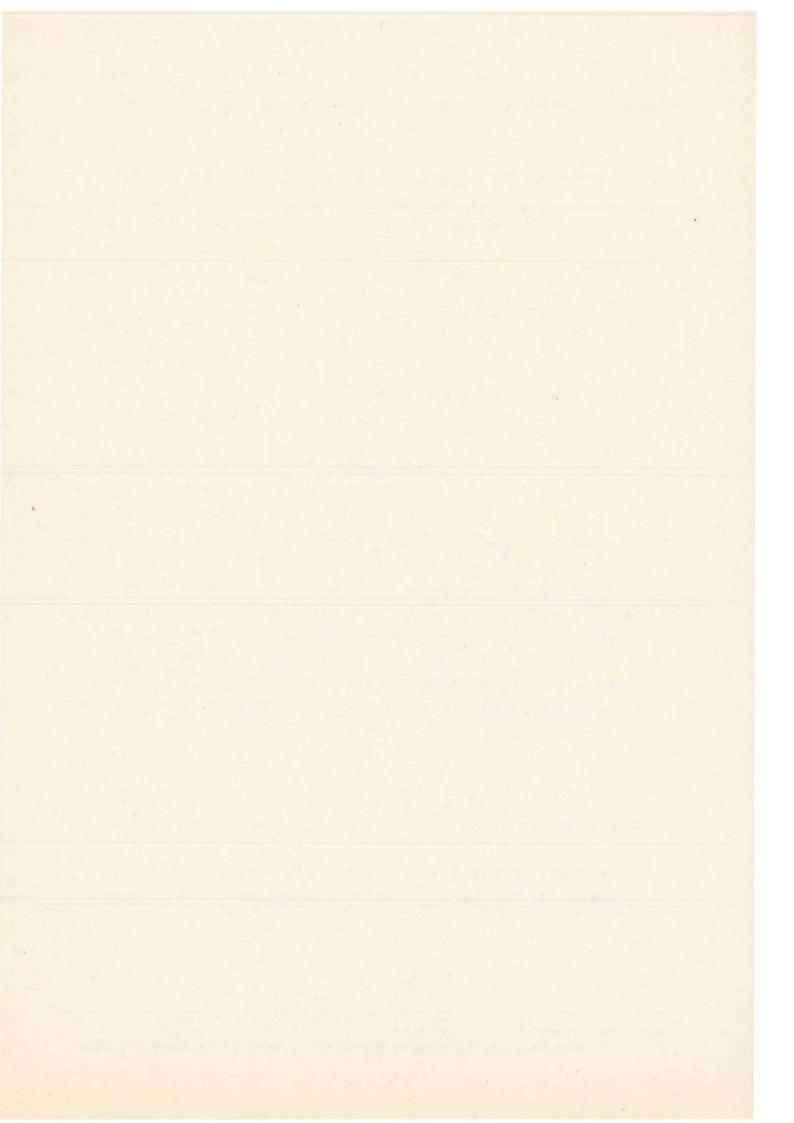
$$Al^{3+}_{limit} = ----- = 1.5 \cdot ----- (5)$$

$$(BC/Al) \qquad (BC/Al)_{crit}$$

L_{min} = min. leaching of base cations [eq/(ha.yr)]

The factor 1.5 derives from the conversion of critical loads

(2)



and base cation concentrations in equivalents to the molar ratio. L_{min} is the minimum leaching of base cations, corresponding to the leaching caused by the residual concentration of base cations that the trees cannot take up. The limiting BC concentration is in the range 10-20 µeq/l; an average of 15 µeq/l has been used:

 $L_{min} = Q.0,015$ (6)

Operationally the H⁺-concentration can be calculated using the gibbsite equation:

$$[H^{+}] = (----)^{1/3}$$
(7
Kgibb

 $K_{gibb} = Gibbsite coefficient = 300 m^6/eq^2$

Accordingly the limiting H⁺-concentration corresponding to a certain Al³⁺-concentration in the soil is calculated from the Al-limiting flux (equation 5), divided by the runoff rate and the gibbsite coefficient:

$$[H^{+}]_{limit} = (-----)^{1/3}$$

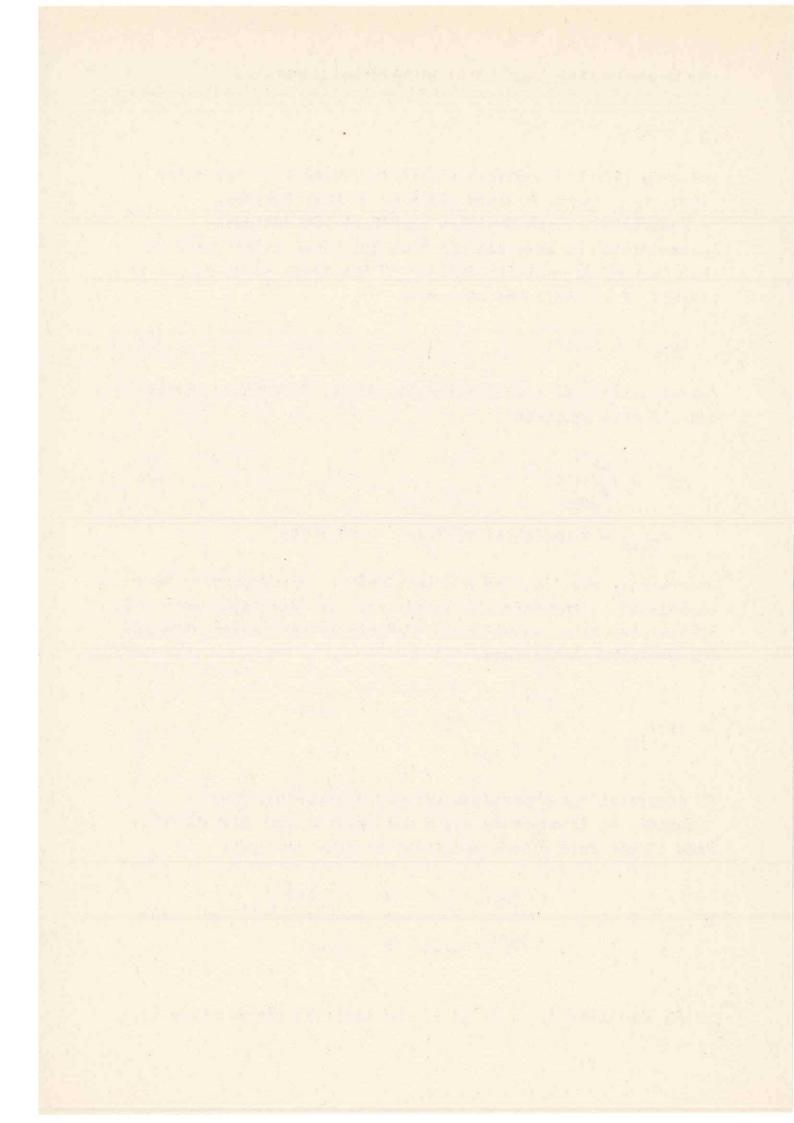
$$Q \cdot K_{gibb}$$
(8)

By inserting the expression for the Al-limiting flux (equation 5) in equation 8 and multiplying with the runoff rate to get from H^+ -concentration to flux, one gets:

 $H^{+}_{limit} = (1.5 . \frac{(1-X_{Na}) . W + D - U - Q.0.015}{(BC/Al)_{crit} . Q . K_{gibb}} (9)$

Using equations 2, 3, 5 and 9, the modified SMB equation for

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the critical loads of actual acidity becomes:

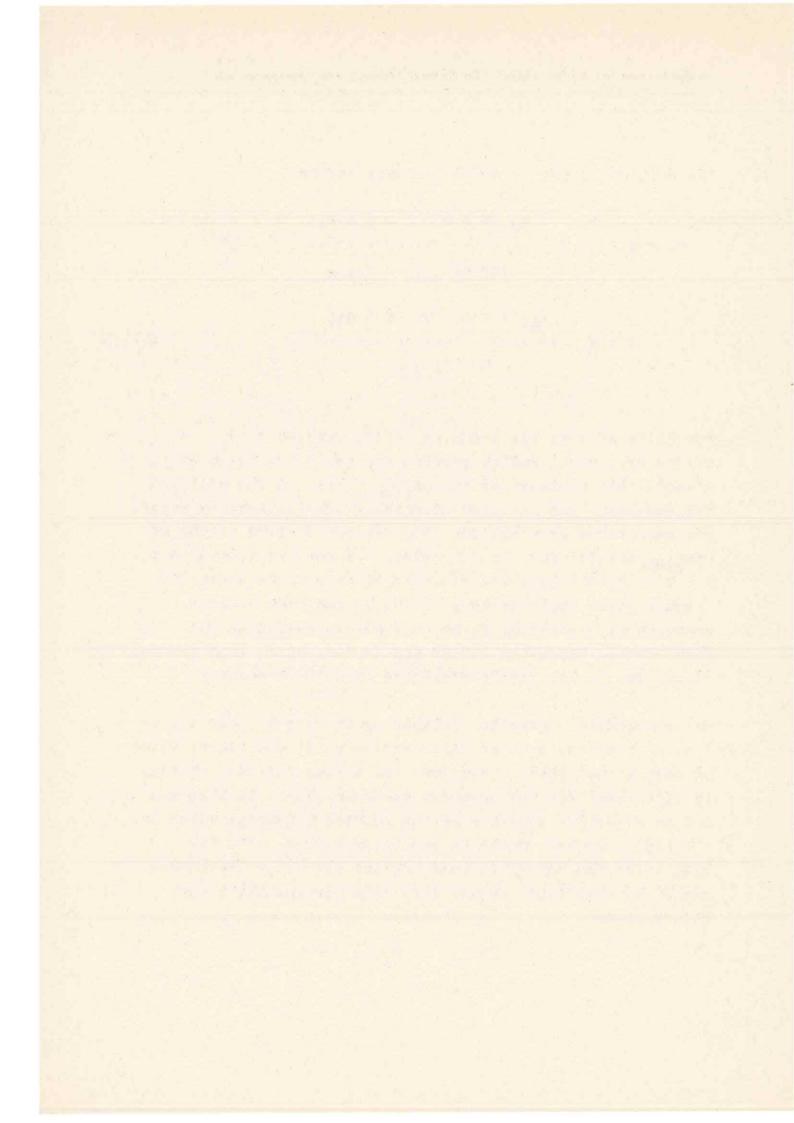
 $CL = W + (1.5 . \frac{x_{BC} . W + D - U - Q.0.015}{(BC/Al)_{crit} \cdot K_{gibb}} \cdot Q^{2/3} + \frac{x_{BC} . W + D - U - Q.0.015}{(BC/Al)_{crit}}$

For units of critical loads in eq/(ha.yr) and Q in $m^3/(ha.yr)$, the gibbsite coefficient takes the value 300 m^6/eq^2 . This corresponds to $-pK_{gibb} = 8.5$. In the modified SMB equation, the dilution of acidity with increasing runoff has been taken into account. Empirically derived values of $-pK_{gibb}$ are 6.5 for the O/A-layer, 7.5 for the E-layer and 8.5 for the B-layer. The modified equation also takes into account, that approximately 20-40% of the base cations produced by weathering is Na, not participating in the BC/Al-ratio. Hence the factor X_{BC} is 0.8, which is a maximum value. The factor varies depending on soil mineralogy.

The simplified expression implies by setting H^+ -leaching to 0,09.Q, that there is no large variation of the runoff rates between sites. This is not true for alpine regions, whereas it is correct for the northern European plain. Furthermore the term (W+D-U) is not constant between different sites in the alpine region. Hence it can be concluded, that the simplified SMB equations were neither valid for the alpine region nor any other region with high precipitation and runoff rates.

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(10)



7

Results

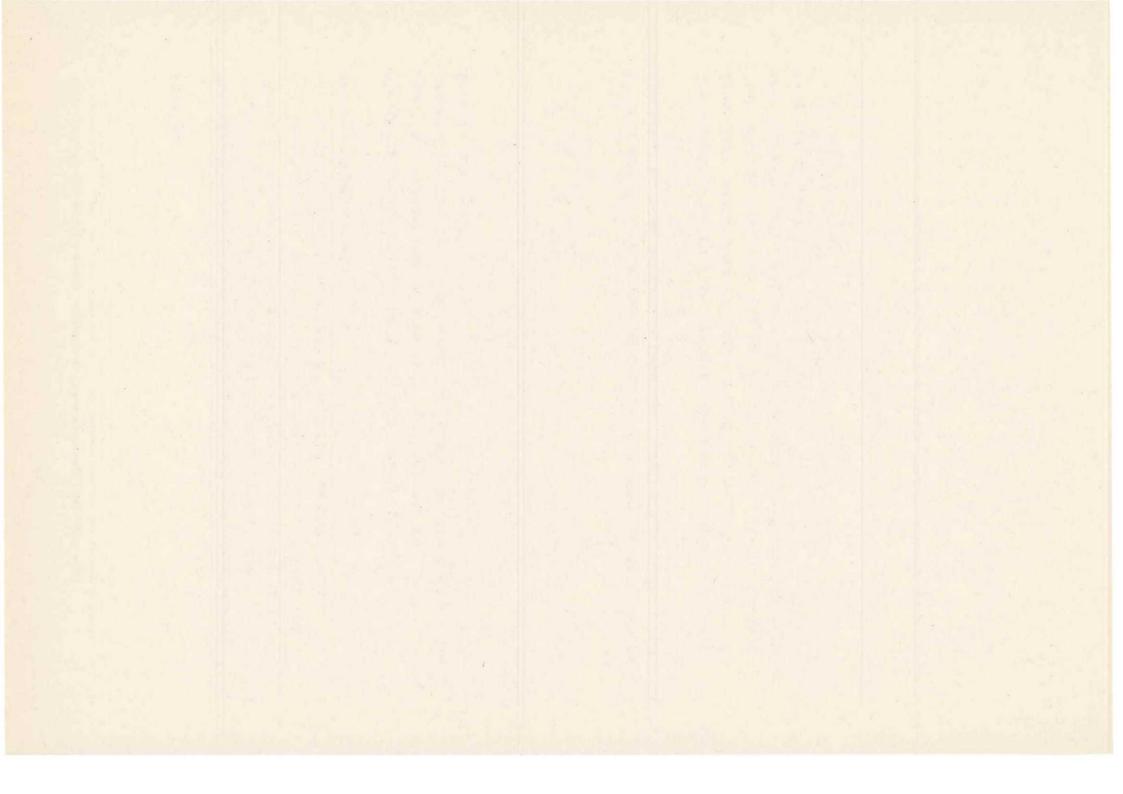
The modified simple mass balance equations for calculating critical loads of actual acidity were applied to the Austrian, German and Swiss datasets for a comparison with the old calculations. The existing databases were used for the calculations; new data were not required, except annual average temperature.

For the Austrian and Swiss datasets additional changes were made. The weathering rate was set proportional to temperature, using an Arrhenius relationship with A=3500. The formula used was:

$$W = W_0 \cdot 10^{-10} \frac{A}{281} \frac{A}{273 + T}$$
(11)

Rooting depth was assumed to decline from 50cm at 1500m above sea level to 10cm at 2500m above sea level.

The figures 1-3 show the Austrian, German and Swiss critical loads maps calculated with the originally proposed simple mass balance equations and the modified ones. It can be seen that the very high critical loads in high precipitation areas have disappeared.



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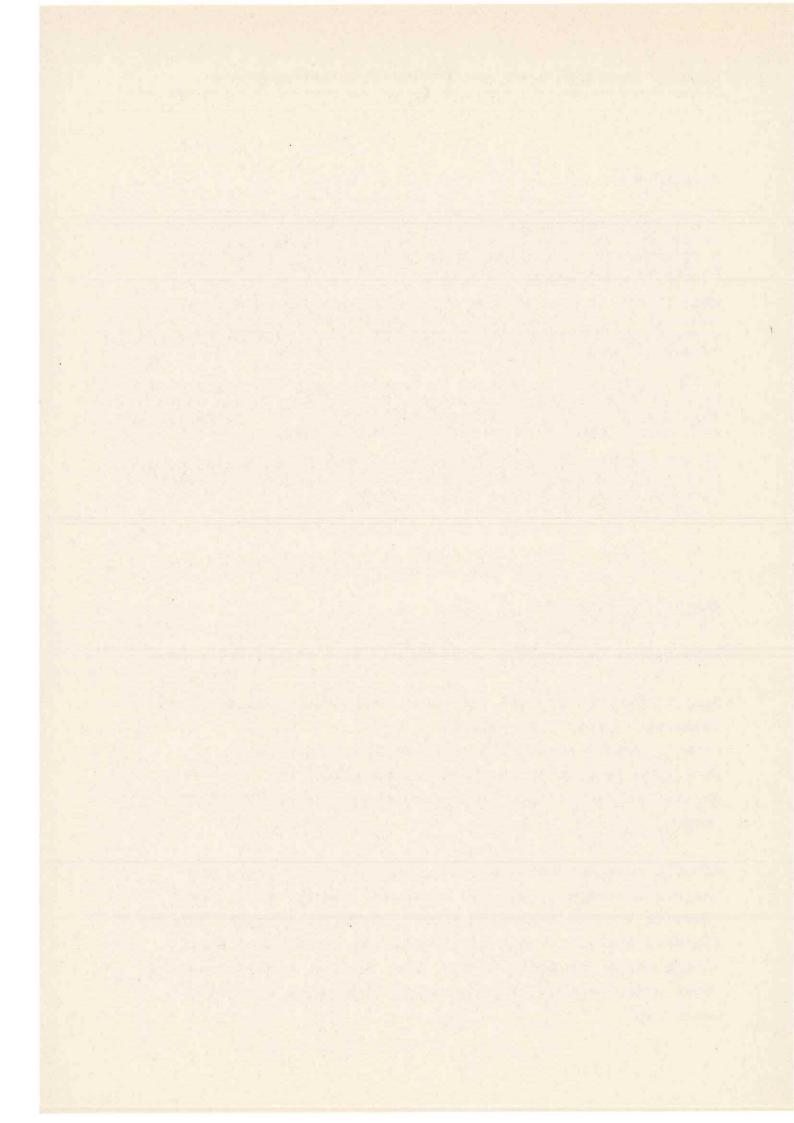
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Acknowledgements:

This report was jointly prepared by (in alphabetical order):

Beat ACHERMANN (Federal Office of Environment, Forests and Landscape, Bern, Switzerland) Helmut HOJESKY (Umweltbundesamt, Vienna, Austria) Beat RIHM (Meteotest AG, Bern, Switzerland) Harald SVERDRUP (Dept. of Chemical Engineering, LTH Lund, Sweden)

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<u>Annex:</u> List of Participants of the Workshop on the Modifications of Austrian, German and Swiss Maps of Critical Loads of Acidity

Dr.Beat ACHERMANN BUWAL Abt.Luftreinhaltung

<u>CH - 3003 BERN</u> Schweiz Tel: +41-31-61 99 78 Fax: +41-31-26 15 46

Dr.Markus AMANN IIASA Dep. TAP Schloßplatz 1 <u>A - 2361 LAXENBURG</u> Österreich Tel: +43-2236-71521-432 Fax: +43-2236-71313

Dr.Ruth BAUMANN UBA Abt. Lufthygiene Spittelauer Lände 5 <u>A - 1090 W I E N</u> Österreich Tel: +43-1-31 30 4-852 Fax: +43-1-31 30 4-400

Dr.Peter BLASER Eidg. Forschungsanstalt f.Wald, Schnee u.Lands.

<u>CH - 8903 BIRMENSDORF</u> Schweiz Tel: +41-1-739-2265 Fax: +41-1-739-2215

Dr.Heinz-Detlef GREGOR UBA I 3.3

Bismarckplatz 1 <u>W - 1000 BERLIN 33</u> Deutschland Tel: +49-30-23-145 846/844 Fax: +49-30-23-156 38 Dipl.Ing.Friedl HERMAN FBVA

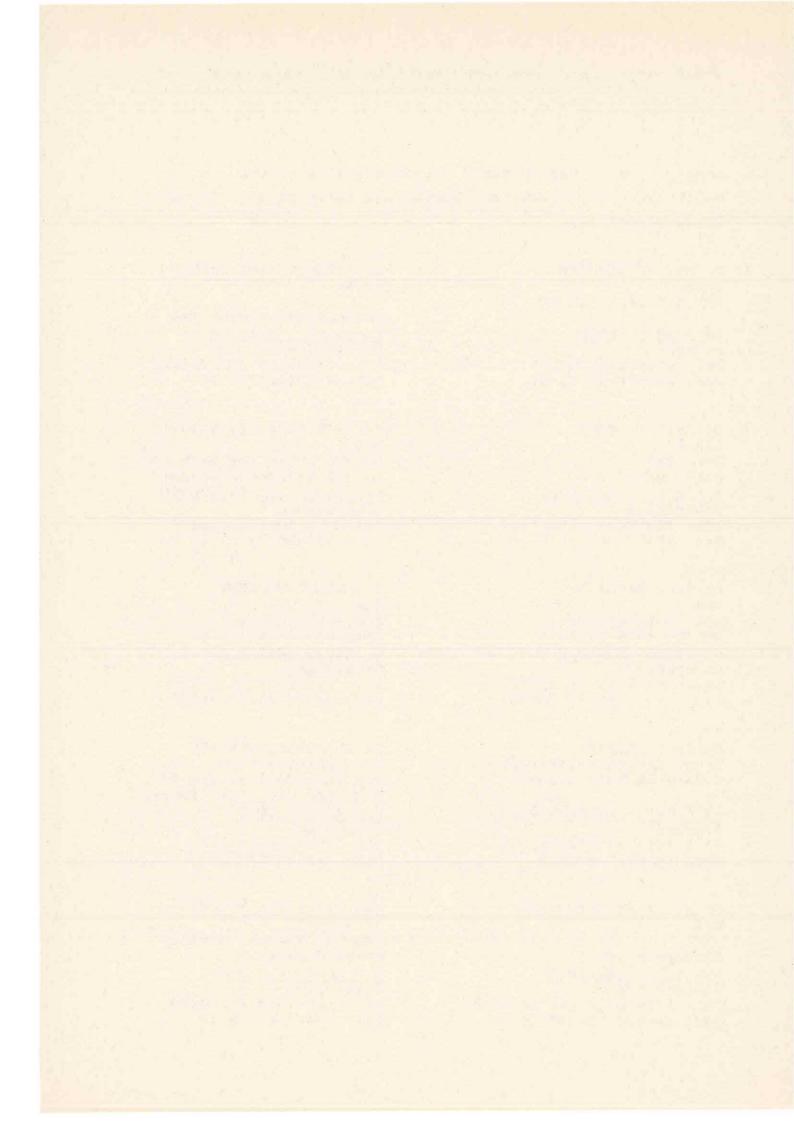
Seckendorff-Gudent Weg 8 <u>A - 1131 W I E N</u> Österreich Tel: +43-1-87 838-105 Fax: +43-1-87 75 907

Dr.Jean-Paul HETTELINGH RIVM Coord.Center for Effects A. van Leeuwenhoeklaan 9 <u>NL - 3720 BA BILTHOVEN</u> Niederlande Tel: +31-30-74 35 29 Fax: +31-30-25 07 40

Dr.Helmut HOJESKY UBA Abt.Lufthygiene Spittelauer Lände 5 <u>A - 1090 W I E N</u> Österreich Tel: +43-1-31 30 4-862 Fax: +43-1-31 30 4-400

Dr.Charlotta JÖNSSON Lund University Dep.Chem.Engineering II P.O. Box 124 Chem. Center <u>S - 22100 LUND</u> Schweden Tel: +46-46-108 274 Fax: +46-46-146 030

Dipl.Ing.Anne KASPER TU Wien Inst.f. Analyt. Chemie Getreidemarkt 9 <u>A - 1060 W I E N</u> Österreich Tel: +43-1-58 801-4979 Fax: +43-1-56 78 13



Dr.Markus KNOFLACHER ÖFZS Abt. Umweltplanung

<u>A - 2444 SEIBERSDORF</u> Österreich Tel: +43-2254-80-2181 Fax: +43-2254(4)-80-2118

Renate KÖBLE Univ. Stuttgart Institut f.Navigation Keplerstr.11 D - 7000 STUTTGART 1 Deutschland Tel: +49-711-121-2166 Fax: +49-711-121-3500

Andreas KOVAR TU Wien Inst.f. Analyt. Chemie Getreidemarkt 9/151 <u>A - 1060 W I E N</u> Österreich Tel: +43-1-58 801-4830 Fax: +43-1-56 78 13

Ir.Hans KROS Winand Staring Centre

P.O.Box 125 <u>NL - 6700 AC WAGENINGEN</u> Niederlande Tel: +31-8370-74 366 Fax: +31-8370-24 812

Dr.Wolfgang LOIBL ÖFZS Abt. Umweltplanung

<u>A - 2444 SEIBERSDORF</u> Österreich Tel: +43-2254-80-2167 Fax: +43-2254-80-2118 Dr.Hans-Dieter NAGEL Inst.f. Ökosystemforsch. Magdalenenstr. 17-19 Postfach 120 <u>O - 1130 BERLIN</u> Deutschland Tel: +37-2-2372-8070 Fax: +37-2-2372-2034

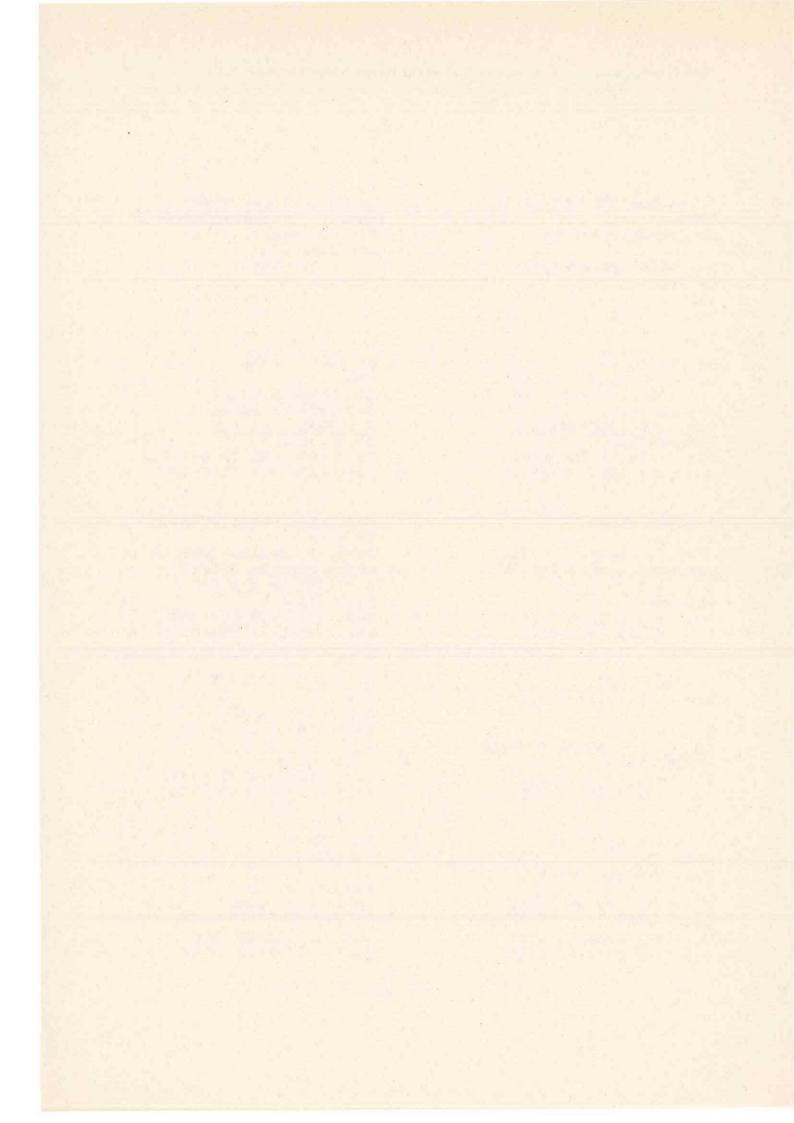
Dr.Horst NOWAK UBA Abt.Umweltplanung Spittelauer Lände 5 <u>A - 1090 W I E N</u> Österreich Tel: +43-1-31 30 4-412 Fax: +43-1-31 30 4-400

Univ.Prof. Hans PUXBAUM TU Wien Inst. f. Analyt. Chemie Getreidemarkt 9/1514 <u>A - 1060 W I E N</u> Österreich Tel: +43-1-58 801-4839 Fax: +43-1-56 78 13

Dr.Klaus RADUNSKY UBA Abt.Lufthygiene Spittelauer Lände 5 <u>A - 1090 W I E N</u> Österreich Tel: +43-1-31 30 4-863 Fax: +43-1-31 30 4-400

Beat RIHM METEOTEST

Fabrikstr. 29 <u>CH - 3012 BERN</u> Schweiz Tel: +41-31-23 74 17 Fax: +41-31-23 42 64



Dipl.Ing. Wolfgang SCHÖPP IIASA Dep. TAP Schloßplatz 1 <u>A - 2361 LAXENBURG</u> Österreich Tel: +43-2236-71 52 1-309 Fax: +43-2236-71 313

Dipl.Ing.Sigrid SCHWARZ UBA Abt.Ökologie Spittelauer Lände 5 <u>A - 1090 W I E N</u> Österreich Tel: +43-1-31 30 4-320Fax: +43-1-31 30 4-400

Dr.Gerhard SMIATEK Univ. Stuttgart Inst. f. Navigation Keplerstr. 11 D - 7000 STUTTGART Deutschland Tel: +49-711-121-3745 Fax: +49-711-121-3500

Dr.Stefan SMIDT FBVA

Seckendorff-Gudent-Weg 8 <u>A - 1131 W I E N</u> Österreich Tel: +43-1-87 838-124 Fax: +43-1-87 75 907

Dr.Christian SMOLINER BM f. Wissenschaft & Forschung, Abt. 23 Freyung 1 A - <u>1010 W I E N</u> Österreich Tel: +43-1-53 120-6353 Fax: +43-1-53 120-6480 Dipl.Ing.Wolfgang SPANGL UBA Abt. Lufthygiene Spittelauer Lände 5 <u>A - 1090 W I E N</u> Österreich Tel: +43-1-31 30 4-861 Fax: +43-1-31 30 4-400

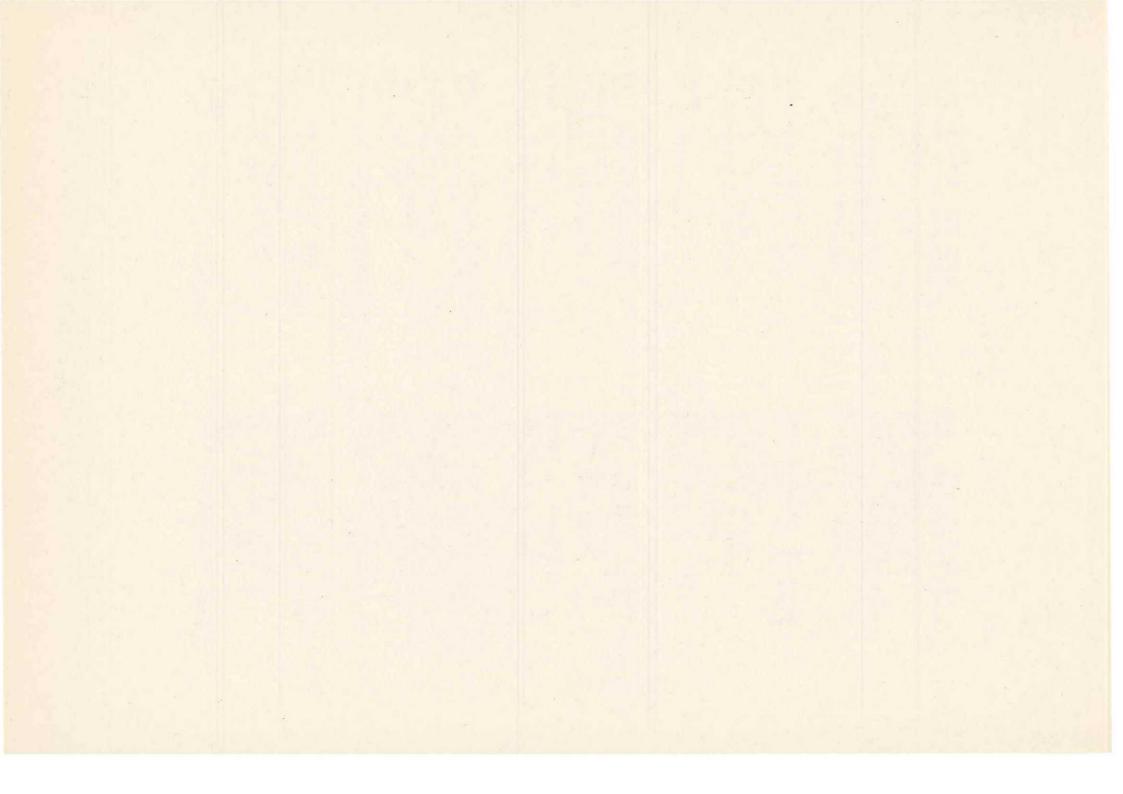
Dr.Harald SVERDRUP Lund University Dep.Chem.Engineering II P.O. Box 124 Chem. Center <u>S - 22100 LUND</u> Schweden Tel: +46-46-108 274 Fax: +46-46-146 030

Dr.Richard VOLZ BUWAL Eidg. Forstdirektion

<u>CH - 3003 BERN</u> Schweiz Tel: +41-31-67 77 86 Fax: +41-31-26 15 46

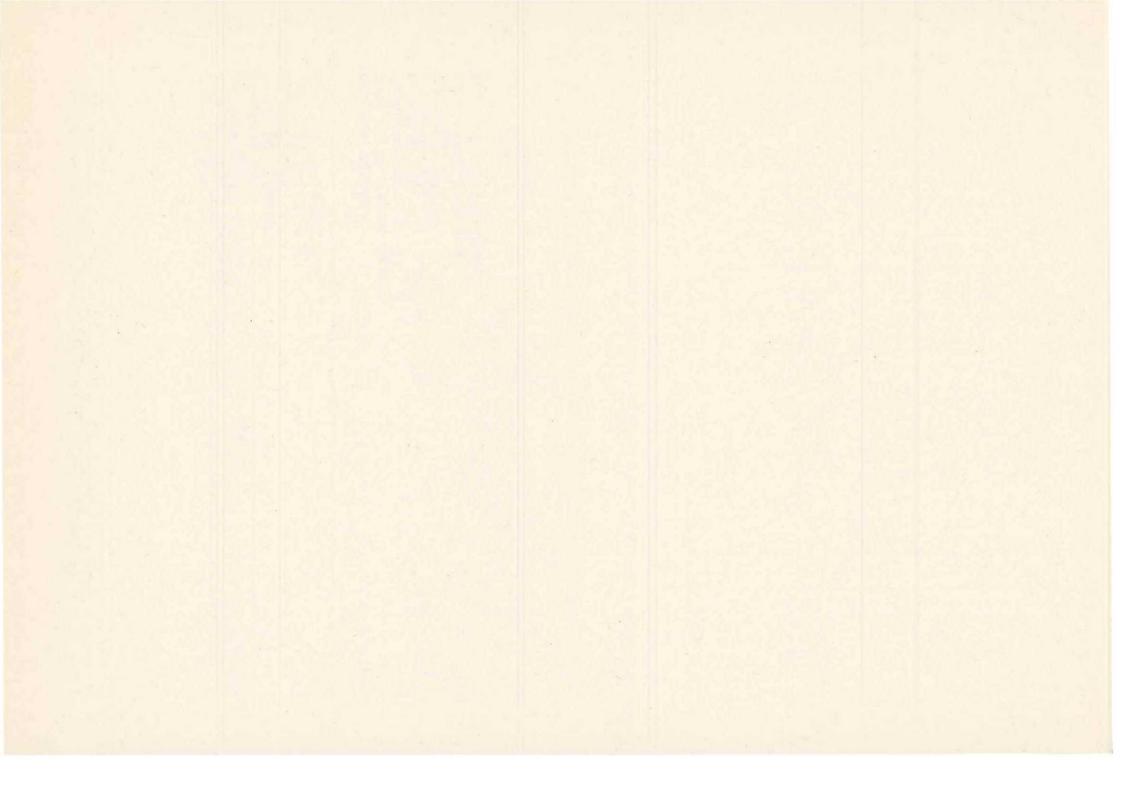
Peter WEISS UBA Abt. Ökologie Spittelauer Lände 5 <u>A - 1090 W I E N</u> Österreich Tel: +43-1-31 30 4-316 Fax: +43-1-31 30 4-400

Dr.Walter WENZEL Univ. f. Bodenkultur Inst. f. Bodenforschung Gregor Mendel-Str. 33 <u>A - 1180 W I E N</u> Österreich Tel: +43-1-34 25 00-306 Fax: +43-1-310 60 27



Beate WERNER Inst.f. Ökosystemforsch. Magdalenenstr. 17-19 Postfach 120 <u>O - 1130 BERLIN</u> Deutschland Tel: +37-2-2372-3987 Fax: +37-2-2372-2034

Dr.Lutz WERNER Inst.f. Ökosystemforsch. Magdalenenstr. 17-19 Postfach 120 <u>O - 1130 BERLIN</u> Deutschland Tel: +37-2-2372-5237 Fax: +37-2-2372-2034



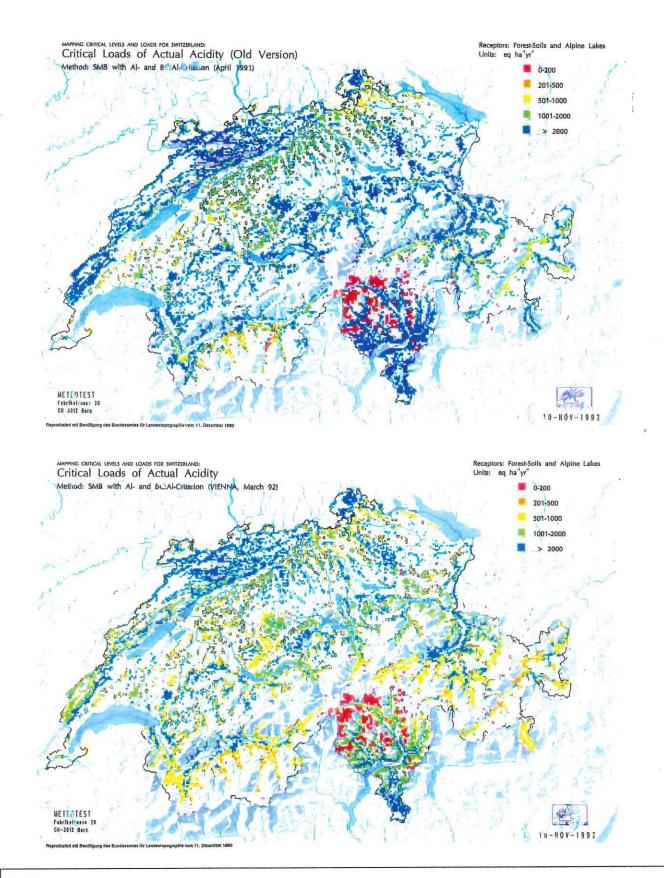


Figure 1: Comparison of the two maps of critical loads of actual acidity for Austria. The map on the top is the old map using the simplified equation; the map on the bottom is the new map using the modified equation. Because of the large bias towards high precipitation – especially in the Alpine regions of Austria – the old map has never been used. The full equation has been applied at an early stage. The only modification done for the new map was the inclusion of the effect of temperature on weathering and the linear decrease of soil rooting depth above 1.500 m a.s.l.

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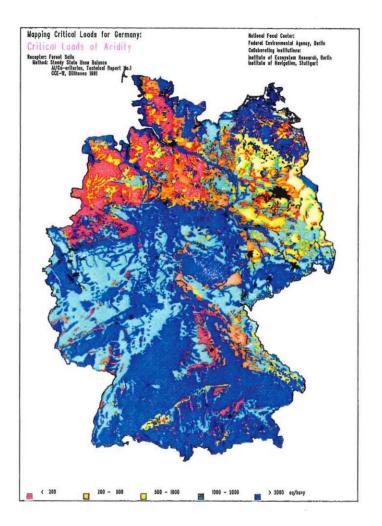
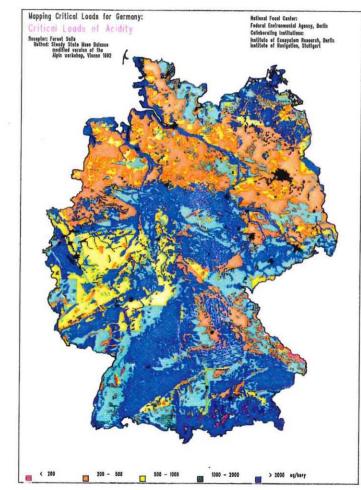
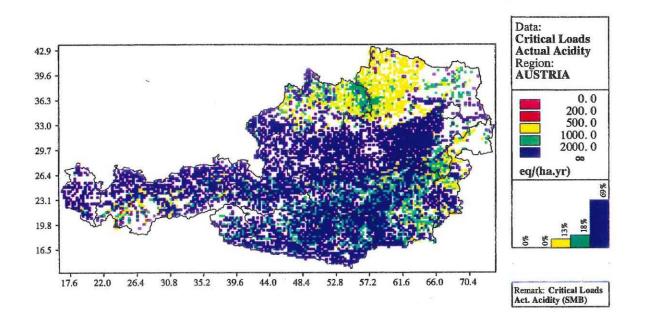


Figure 2: Comparison of the two maps of critical loads of actual acidity for Germany. The map on the top is the old map using the simplified equation; the map on the bottom is the new map using the modified equation. The old map shows some ambiguities (especially in the Harz and Fichtel mountains). The new map shows a more diverse picture; areas with forest damage appear as sensitive. The bias towards high precipitation is not longer present.





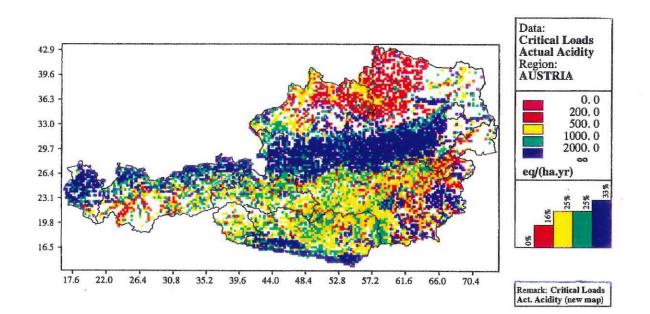


Figure 3: Comparison of the two maps of critical loads of actual acidity for Switzerland. The map on the top is the old map using the simplified equation; the map on the bottom is the new map using the modified equation. It can be seen in the old map, that large parts of Switzerland remain very acid tolerant. Even areas with well documented forest damage such as Jura, Lower Alps, Alps and Ticino remain unsensitive. With the new map there is much better correspondence between the spatial distribution of critical loads and areas known to be sensitive to acidic deposition.